# Gas transfer velocities at water and rough-surface interfaces Gabriel Katul<sup>1,2,3 &</sup> Heping Liu<sup>4</sup>



**Ross Barnett Reservoir**, Mississippi, USA. *Photo credit*: Heping Liu

A view of the large 20foot-diameter lysimeter at Davis planted to beans in 1968.



University of California, Davis, CA **weighing lysimeter**. From *Pruitt et al. (1972)* 

#### Presentation for the TERENO Workshop, Garmisch, Germany (26.9.2017)

<sup>1</sup>Nicholas School of the Environment and <sup>2</sup>Department of Civil and Environmental Engineering, Duke University, Durham, North Carolina, USA (email: gaby@duke.edu)

<sup>3</sup> MICMoR Visiting Scientist, KIT/IMK-IFU, Garmisch-Partenkirchen, Germany

<sup>4</sup>Department of Civil and Environmental Engineering, Washington State University, Pullman, Washington State, USA

## Introduction

 The significance of air-water gas exchange on ecosystem and climate-related processes (e.g. CO<sub>2</sub> exchange with open water surfaces) is not in dispute. Likewise for gas exchange between a turbulent atmosphere and a rough surface (e.g. evaporation from soil to atmosphere, nitrous oxide emissions,..).

 However, the mechanisms describing the *efficiency* of gas transport by turbulence at such interfaces remains multifaceted and subject to active research.

## Introduction

There is a '*renaissance*' in mass transfer studies across *interfaces* because of rapid advancements in <u>measurements</u> (e.g. remote sensing using high resolution IR cameras<sup>1,2</sup>, fiber optic cables<sup>3</sup>, stable isotopic techniques<sup>4</sup>), and <u>simulations</u> (LES and DNS<sup>5</sup>):

#### Examples:

- <sup>1</sup>Bare soil evaporation: Haghighi, E., and D. Or (2013, 2015).
- <sup>2</sup>Air-sea exchange of  $CO_2$  and sparingly soluble gases: Garbe et al. (2004).
- <sup>3</sup>Crop evapotranspiration by Bowen ratio: Euser et al. (2014).
- <sup>4</sup>Sediment-water interface: Merlivat, L., and M. Coantic (1975).
- <sup>5</sup>Ocean-atmosphere gas exchange: Fredriksson et al. (2016).

## Introduction

- Instead of dealing with fluxes and concentration differences, it is preferable to deal with 'gas transfer velocity' when quantifying turbulent transport efficiency.
- Flux:  $F = \mathbf{k} \Delta C$ ;  $\Delta C = C_s C_b$

$$\boldsymbol{k} = \frac{F}{\Delta C}.$$

• The *efficiency* of gas transport by turbulence at such interfaces must account for eddies.

## **Review of k-Formulations**

- Common k formulations that account for turbulent eddies are based on surface renewal schemes subject to assumptions about contact durations at the interface.
- For an air-water interface, the water-side gas transfer velocity(Lamont and Scott, 1970)  $k = \alpha Sc^{-n} (\epsilon v)^{1/4}$
- For gas transfer from rough surfaces into a turbulent atmosphere (*Brutsaert, 1965*)

$$k = A \sqrt{D_m} u_*^{3/4} (\kappa_v \nu z_o)^{-1/4}$$

# Objective

- Show that these formulations inherit their universal character from Kolmogorov's 1941 inertial subrange scaling adjusted by viscous cutoff.
- The proposed derivation explains why
- (i) gas transfer models are insensitive to renewal time distributions, and
- (ii) the similarity constants (α or A) are independent from the specifics of the interfacial system.

# Definitions and review of air-water gas transfer theories



Eibsee Lake, Bavaria, Germany: Photo Credit - Alessandra Trompeo

## Gas Transfer Velocity at an Air-Water Interface

### Transfer Velocity

• Flux:  $F = \mathbf{k} \Delta C$ ;

• Concentration Difference:

 $\Delta C = C_s - C_b;$ 

 C<sub>s</sub> surface concentration determined from gas phase measurements and Henry's Law (or the Ostwald solubility coefficient).



Gas transfer velocity models: What is done in practice

• Early formulations for k - still in use in climate models - are based on mean velocity U.

Based on Cole and Caraco's (1998) review:

$$k = a + bU^{c}$$

- <u>a:</u> 0 2.07
- <u>b:</u> 0.2 0.75
- <u>c:</u> 1-2

Mean wind speed at a reference height (often 10 m)

## Global maps of k for inland water

• From Raymond et al. (2013) – based on  $k = a + bU^c$ 



## Formulation of gas transfer velocity at air-water interface

 A detailed formulation for k is based on surface renewal/micro-eddy theory (Lamont and Scott, 1970).

• 
$$k = \alpha Sc^{-n} (\epsilon \mathbf{v})^{1/4};$$

An Eddy Cell Model of Mass Transfer into the Surface of a Turbulent Liquid

JOHN C. LAMONT and D. S. SCOTT

University of Waterloo, Waterloo, Ontario, Canada

- $\alpha = 0.4$  (but 0.2-0.6 reported range)
- Sc = Molecular Schmidt Number =  $\nu/D_m$  (>>1 for water)
- v = Kinematic viscosity of water,  $D_m$  = molecular diffusion of C
- $\epsilon$  = Water-side mean turbulent kinetic energy (**TKE**) dissipation rate
- n = 1/2 (but 2/3 is for smooth wall).

## The meaning of $(\epsilon \nu)^{1/4}$ **Kolmogorov scales for small eddies**

 $\eta = \left(\frac{\nu^3}{\epsilon}\right)^{1/4}$  is the Kolmogorov length scale, smallest scale of turbulence where turbulent kinetic energy is dissipated by the action of viscosity.

$$\tau_k = \sqrt{\left(\frac{\nu}{\epsilon}\right)}$$
 is the Kolmogorov time scale  
 $(\epsilon \nu)^{1/4} =$  is the Kolmogorov velocity scale

*Microscale Reynolds* number:  $Re_K = [\eta \ (\epsilon 
u)^{1/4}]/
u = 1$ . Photo: John Collier - Copied from johnbyrne.fireflyinternet.co.uk. Cropped photo of a painting of Osborne Reynolds painted in 1904 by John Collier.



## Experiments



Click Here

Full Article

# Review of gas transfer theories from rough surfaces into a turbulent atmosphere



Image from https://ohiocountrylife.files.wordpress.com/2015/03/bare-soil.jpg



JOURNAL OF GEOPHYSICAL RESEARCH

Vol. 70, No. 20

A Model for Evaporation as a Molecular Diffusion Process into a Turbulent Atmosphere

WILFRIED BRUTSAERT

Department of Hydraulics and Hydraulic Engineering School of Civil Engineering, Cornell University, Ithaca, New York

Abstract. Evaporation is analyzed as a phenomenon of molecular diffusion from a large rough surface into random-lived internal scale turbulent eddies. The model is based on the principles of similarity and stochastic renewal. It is shown that evaporation is proportional to the 0.75 power of the friction velocity and to the square root of the molecular diffusivity. This is in agreement with the available experimental data.



## Brutsaert (1965)

A simplified view of Brutsaert's renewal model:

• Evaporation at any time *t* is:

$$E(t) = -D_m \frac{\partial C(t)}{\partial z} \bigg|_{z=0} \approx D_m \left(\frac{\Delta C}{\delta(\tau)}\right) \sim D_m \frac{\Delta C}{\sqrt{D_m \tau}}$$

 $\Delta C = (C_s - C_b)$  assumed time independent.

$$\frac{E(t)}{\Delta C} = k(\tau) \sim \sqrt{\frac{D_m}{\tau}}; \tau > 0$$



mean contact duration  $\propto$  <u>Kolmogorov time scale</u>

$$k = \sqrt{A\pi D_m} \left(\frac{\nu}{\epsilon}\right)^{-1/4} = \sqrt{A\pi}Sc^{-1/2} (\epsilon\nu)^{1/4}$$
  

$$\epsilon = u_*^2 \left(\frac{dU}{dz}\right) = u_*^2 \left(\frac{u_*}{\kappa_{\nu} z_o}\right)$$
Turbulent kinetic energy  
budget at  $z_o$  - Dissipation  
mechanical production  

$$E = \sqrt{A\pi D_m} u_*^{3/4} (\kappa_{\nu} \nu z_o)^{-1/4} \Delta C$$

\*Result not sensitive to assumed exponential pdf of  $\tau$  (Katul and Liu, 2017a)

## Brutsaert (1965)

• Average 
$$\delta = -\int_0^\infty a e^{-a\tau} \sqrt{D_m \tau} d\tau = \frac{1}{2} \sqrt{\frac{\pi D_m}{a}}$$
.  
• Recall that  $a = \frac{A}{\sqrt{\left(\frac{\nu}{\epsilon}\right)}}; \delta = \frac{1}{2} \sqrt{\frac{\pi \nu}{a \, Sc}}$ .

δ ~ Sc<sup>-1/2</sup> η. Hence, the most efficient eddy size appears to be the Batchelor length scale (irrespective of the Schmidt number).



## **THEORY: A Structure Function Perspective**

Image: http://www.russia-ic.com/education\_science/gems/817/1345938588000#.WcI9aNEpBPY

**Transfer Velocity:** 

$$\boldsymbol{k} \approx \sqrt{(w(x+r) - w(x))^2} = \sqrt{D_{ww}(r)}$$



Assumption 1:  $k = \sqrt{C_o \epsilon^{2/3} r^{2/3}}$  K41 scaling for  $D_{ww}(r)$ 

Assumption 2:  $r \sim Sc^{-1/2} \eta$  (Hondzo, 1998)

For many gases in the atmosphere -  $Sc \approx 1$   $r \sim \eta$ ; where  $\eta = \left(\frac{\nu^3}{\epsilon}\right)^{1/4}$  Kolmogorov microscale



## Structure function model for

## air-water gas exchange

From Wanninkhof et al. (2009)			
Gas	Mol. weight (g mol <sup>-1</sup> )	D (10 <sup>-5</sup> cm <sup>2</sup> s <sup>-1</sup> )	Sc
<sup>3</sup> He	3	7.29	144
He	4	6.36	165
CH <sub>4</sub>	16	1.55	677
Ne	20	3.33	315
$N_2$	28	1.57	670
O <sub>2</sub>	32	1.78	589
Ar	40	1.82	576
CO <sub>2</sub>	44	1.59	660
N <sub>2</sub> O	44	1.5	698
(CH <sub>3</sub> ) <sub>2</sub> S	62	1.14	918
Kr	84	1.51	694
CCl <sub>2</sub> F <sub>2</sub>	121	0.86	1219
Xe	131	1.19	880
CCl <sub>3</sub> F	137	0.94	1120
SF <sub>6</sub>	146	1.05	992
CCl <sub>4</sub>	154	0.82	1286
CCl <sub>2</sub> FCClF <sub>2</sub>	187	0.68	1544
Rn	222	1.07	980
Heat	_	1.75	6

 $Sc \gg 1$ Water molecules are tightly packed compared to air molecules, where Sc is near *unity*.

## **Air-Water**



## **A Structure Function Perspective**

A solution to approximated KH equation (Katul et al., 2015):

$$\frac{D_{ww}(r)}{C_o(\epsilon r)^{2/3}} = 1 - \begin{bmatrix} \frac{1}{\xi} Daw_F(\xi) \\ \frac{1}{\xi} Daw_F(\xi) \end{bmatrix}$$
Viscous corrections to K41 scaling  
$$\xi = \theta(r/\eta)^{2/3}; \ \theta = \sqrt{\frac{1}{10C_o}}.$$
$$Daw_F(\xi) \approx \xi - \frac{2}{3}\xi^3 + \cdots \text{ is the Dawson}$$
function.

## Recovery of $\alpha$ value

• Insert into structure function leads to:

$$k = \sqrt{\frac{2}{15}} Sc^{-1/2} (\epsilon \nu)^{1/4}.$$
  
Here,  $\alpha = \sqrt{\frac{2}{15}} = 0.36$ 

Recall Zappa et al. (2007):  $\alpha = 0.41$ 

This completes the sought-after result and the study objective. A final *independent* check – recovery of surface divergence theories.

## **Recovery of Surface Divergence Methods**

From Banerjee et al. (2004):

DNS & Measurements  

$$0.57 \cdot 0.71$$
  
 $k = c_s Sc^{-1/2} \sqrt{\nu \Lambda_o}$ ;  $\Lambda_o = \sqrt{\left(\frac{\partial w}{\partial z}\right)^2}$ .  
Start with  $k = \sqrt{\frac{2}{15}} Sc^{-1/2} (\epsilon \nu)^{1/4}$  (i.e. Structure function)  
Recall that  $\epsilon = 15\nu \left(\frac{\partial w}{\partial z}\right)^2$  (in locally isotropic turbulence)

$$k = \frac{\sqrt{2}}{15^{1/4}} Sc^{-1/2} (\mathbf{v}\Lambda_o)^{1/2}.$$

## Conclusion

- $k = \frac{F}{\Delta C}$  (analogous to a macroscopic resistance law).
- $k = [(w(x + r) w(x))^2]^{1/2}$  structure function describes cumulative energy content at scale r (e.g. analogous to micro-state energy content). Most *effective eddy size r* contributing to k is linked to the Kolmogorov microscale (or Batchelor scale).
- The equality of these two expressions may be viewed as analogous to a *'fluctuation-dissipation'* result for gas transfer velocity across interfaces.

# Thank you

All models are wrong, but some are useful (George Box). From https://en.wikiquote.org/wiki/George\_E.\_P.\_Box



Acknowledgements

**E. Bou-Zeid**, **M. Novack**, and **W. Brutsaert** for the discussions and insights on this topic.

MICMoR Visiting Fellows program, KIT/IMK-IFU – Garmisch, Germany

National Science Foundation: NSF-EAR-1344703 and NSF-DGE1068871, NSF-AGS-1112938

Department of Energy: DE-SC0006967 and DE-SC0011461

## References

- Banerjee et al. 2004. Surface divergence models for scalar exchange between turbulent streams. *International Journal of Multiphase Flow, 30*(7), 963-977.
- Brutsaert, W., 1965. A model for evaporation as a molecular diffusion process into a turbulent atmosphere. J. Geophys. Res., 70(20), 5017-5024.
- Cole, J.J. and Caraco, N.F., 1998. Atmospheric exchange of carbon dioxide in a low-wind oligotrophic lake measured by the addition of SF6. *Limn. and Ocean.*, 43(4), 647-656.
- Euser et al., 2014, A new method to measure Bowen ratios using high-resolution vertical dry and wet bulb temperature profiles, *Hydrol. Earth Syst. Sci.*, 18, 2021-2032.
- Fredriksson et al. 2016, An evaluation of gas transfer velocity parameterizations during natural convection using DNS, *J. Geophys. Res. Oceans*, 121, 1400–1423, doi:10.1002/2015JC011112.
- Garbe et al. 2004. A surface renewal model to analyze infrared image sequences of the ocean surface for the study of air-sea heat and gas exchange. J. Geophys. Res.: Oceans, 109(C8).
- Haghighi, E., and D. Or., 2013. Evaporation from porous surfaces into turbulent airflows: Coupling eddy characteristics with pore scale vapor diffusion. *Water Resources Research* 49,12: 8432-8442.
- Haghighi, E. and Or, D., 2015. Evaporation from wavy porous surfaces into turbulent airflows. *Transport in Porous Media*, 110(2), 225-250.
- Hondzo, M., 1998. Dissolved oxygen transfer at the sediment-water interface in a turbulent flow. *Water Resources Research*, *34*(12), 3525-3533.
- Katul, G. and Liu, H., 2017a. A Kolmogorov-Brutsaert structure function model for evaporation into a turbulent atmosphere. *Water Resources Research*, 53(5), 3635-3644.
- Katul, G. and Liu, H., 2017b. Multiple mechanisms generate a universal scaling with dissipation for the air-water gas transfer velocity. *Geophysical Research Letters*, 44(4), 1892-1898.
- Katul et al. 2015. Bottlenecks in turbulent kinetic energy spectra predicted from structure function inflections using the von Kármán-Howarth equation. *Physical Review E*, *92*(3), 033009.
- Kolmogorov, A.N., 1941. The local structure of turbulence in incompressible viscous fluid for very large Reynolds numbers. In *Dokl. Akad. Nauk SSSR*, 30, 4, 299-303.
- Lamont, J.C. and Scott, D.S., 1970. An eddy cell model of mass transfer into the surface of a turbulent liquid. *AIChE Journal*, *16*(4), 513-519.
- Merlivat, L. and Coantic, M., 1975. Study of mass transfer at the air-water interface by an isotopic method. J. Geophys. Res., 80(24), 3455-3464.
- Pruitt et al. 1972, Water use by crops as affected by climate and plant factors, *California Agriculture*, October, 10-14.
- Raymond et al, 2013. Global carbon dioxide emissions from inland waters. *Nature*, *503*(7476), 355.
- Takagaki et al. 2016. Effect of Schmidt number on mass transfer across a sheared gas-liquid interface in a wind-driven turbulence. *Scientific reports, 6,* doi:10.1038/srep37059.
- Wang et al. 2015. On the coefficients of small eddy and surface divergence models for the air-water gas transfer velocity. J. Geophys. Research: Oceans, 120(3), 2129-2146.
- Wanninkhof et al. 2009. Advances in quantifying air-sea gas exchange and environmental forcing, *Annual Review of Marine Science*, 213-244, doi: 10.1146/annurev.marine.010908.163742
- Zappa et al. 2007. Environmental turbulent mixing controls on air-water gas exchange in marine and aquatic systems. *Geophysical Research Letters*, 34(10).

## **EXTRA SLIDES FOR QUESTIONS**

# Schmidt Number Effect $k = \alpha Sc^{-n} (\epsilon \nu)^{1/4}$



From: Takagaki et al. 2016

Evaporation derived from an interfacial theory and surface renewal (Brutsaert, 1965)



## Add Viscous Cutoff Correction to K41

to viscous ranges 10<sup>0</sup>  $w_{ww}(r)/C_o(\varepsilon r)^{2/2}$ K41 10<sup>-2|</sup>  $10^{0}$ r/ŋ

Cross-over from inertial

Basic result

$$k \sim (\epsilon \boldsymbol{\nu})^{1/4}$$

Unaltered. Only the constant **A** needs adjustment.

(see Katul and Liu, 2017a)

## **THEORY: A Structure Function Perspective**

ONLY Kolmogorov Scaling: 
$$k = \sqrt{[C_o \epsilon^{2/3} r^{2/3}]};$$
  
 $r = Batchelor Scale l_B = S_c^{-1/2} \eta;$  where  $\eta = \left(\frac{\nu^3}{\epsilon}\right)^{1/4}$ .  
 $k = \sqrt{C_o} Sc^{-1/6} (\nu \epsilon)^{1/4}$ 

Notes:  $(\epsilon v)^{1/4}$  = Kolmogorov velocity K41 scaling leads to correct  $(v\epsilon)^{1/4}$  but wrong Schmidt number exponent

(see Katul and Liu, 2017b).



## Inertial-to-Viscous Crossover

 Model reproduces the 'spectral bottleneck' when physical space is converted to spectral



## Typical profiles of mean concentration

- The shape of the mean concentration profile over time scales much longer than turbulent mixing are 'preserved'.
- This suggests that

$$\Delta C = C_s - C_b$$
 is a

reasonable descriptor of the driving force for mass movement.

50 100 150 200 250 300 0 7 Aug 6 Jul 2 15 Jun Depth (m) 4 6 8 10

SF<sub>6</sub> (pmol liter-1)

From Cole and Caroco (1998)

#### **Transfer Across the Air-Sea Interface**

Christoph S. Garbe, Anna Rutgersson, Jacqueline Boutin, Gerrit de Leeuw, Bruno Delille, Christopher W. Fairall, Nicolas Gruber, Jeffrey Hare, David T. Ho, Martin T. Johnson, Philip D. Nightingale, Heidi Pettersson, Jacek Piskozub, Erik Sahlée, Wu-ting Tsai, Brian Ward, David K. Woolf, and Christopher J. Zappa

#### Peter S. Liss Martin T. Johnson *Editors*

Springer Earth System Sciences

#### Ocean-Atmosphere Interactions of Gases and Particles

